# Moment Tensor Inversion 

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## Introduction

Now, we can compute synthetic seismograms that are comparable with observed seismograms.

The seismic waveforms contain the information of the focal mechanism and seismic slip area.

If we assume earth structure, we can calculate green's function and estimate moment tensor components and location of centroid using waveform inversion scheme.

## Forward modeling and Inversion



To estimate source model, we often apply two method.
Forward Modeling (try and error)
Input: source model
Output: synthetic waveform
Inversion
Input: observation data
Output: source model

## Linear Inversion

Observation Equation

$$
\mathbf{y}=\mathbf{f}(\mathbf{x}) \quad \mathbf{y} \text { : observation; } \mathbf{x}: \text { model }
$$

Linear case

$$
\mathbf{y}=\mathbf{A} \mathbf{x} \quad \mathbf{A}: \text { Kernel Matrix }
$$

And then,

$$
\hat{\mathbf{x}}=\left(\mathbf{A}^{T} \mathbf{A}\right)^{-1} \mathbf{A}^{T} \mathbf{y}
$$

We can get model information from observation data.

## Double Couple Model



The double couple model has two pairs of single couple.
This model don't have total face and torque. Thus this model have net balanced the moment.

## Intrinsic problem

Any time, we can not choose the actual fault plane of two nodal planes with point source. If we want to determine one fault plane, we need to refer to aftershock distribution or tectonic setting.

## Seismic Moment tensor

The moment tensor consists of nine single-couples force in the localsource Cartesian coordinate system.

The explosion source described can be modeled by the sum of the three dipole terms, $M_{11}+M_{22}+$ $M_{33}$, with each having equal moment.

The seismic moment tensor is always symmetric.


## Moment tensor and fault motion

Moment tensor and crack

$$
M_{i j}=\left[\lambda\left(\sum_{k=1}^{3} v_{k} n_{k}\right) \delta_{i j}+\mu\left(v_{i} n_{j}+v_{j} n_{i}\right)\right] D S
$$

Moment tensor and dislication

$$
M_{i j}=\left(v_{i} n_{j}+v_{j} n_{i}\right) \mu D S=\left(v_{i} n_{j}+v_{j} n_{i}\right) M_{0}
$$



## Moment Tensor and Radiation Pattern

Moment tensor and Far-field term

$$
U_{c}(r, \gamma, t)=\frac{R_{c}}{4 \pi \rho c^{3}} \frac{1}{r} Q(t) * \dot{M}_{0}\left(t-\frac{r}{c}\right) \quad \text { Q-effect }_{\text {Moment-rate function }}
$$

## Radiation Pattern

$$
R_{C}=\sum_{i=1}^{3} \sum_{j=1}^{3} m_{i j} \gamma_{i} e_{C j}
$$

## P-wave Radiation Pattern

$$
R_{P}=2\left(\sum_{i=1}^{3} v_{i} \gamma_{i}\right)\left(\sum_{j=1}^{3} n_{j} \gamma_{j}\right)
$$



## Radiation Pattern in Nuclear Weapons Testing

Moment tensor components

$$
m_{i j}=\delta_{i j} \quad i, j=1,2,3
$$

P-wave Radiation Pattern

$$
R_{P}=\sum_{i=1}^{3} \sum_{j=1}^{3} \delta_{i j} \gamma_{i} e_{P j}=\sum_{i=1}^{3} \sum_{j=1}^{3} \delta_{i j} \gamma_{i} \gamma_{j}=\sum_{i=1}^{3} \gamma_{i} \gamma_{i}=1
$$

S-wave Radiation Pattern

$$
R_{S}=\sum_{i=1}^{3} \sum_{j=1}^{3} \delta_{i j} \gamma_{i} e_{S j}=\sum_{i=1}^{3} \gamma_{i} e_{S i}=0
$$

## Basis Moment Tensor

We can treat source and propagation process as linear operators, and the moment tensor can describe doublecouple at any time.
It is possible to construct observed waveform by summing the weighted the green's functions for each basis moment tensor.

We assume only double-couple model, the number of independent components of moment tensor is five.


Kikuchi and Kanamori (I99I, BSSA)

## Green's Function for each basis moment tensor

JB structure model (Moho $=33 \mathrm{~km}$ )
Station KBS (Az.=350, Del. =61)
Depth $=0.5$ km



## Green's Function for each basis moment tensor

JB structure model (Moho $=33 \mathrm{~km}$ )
Station KBS (Az.=350, Del. =61)
Depth = 10 km


## Green's Function for each basis moment tensor

Station KBS (Az.=350, Del. =61)
Depth $=\mathbf{2 0}$ km



## Moment Tensor Inversion

Observed seismic waveform of $c$ component at a station $j$ due to seismic moment release in a source volume $V$

$$
u_{c j}(t)=\sum_{q=1}^{6} \iiint_{\nu} \tilde{G}_{c j q}(t, \xi) * \tilde{M}_{q}(t, \boldsymbol{\xi}) d \xi+e_{c j}^{\prime}(t)
$$

Obs. Waveform Green's function Moment-rate function
Assumption 1: Point source model, in which we assume the seismic waveform to be radiated from one point.

$$
u_{c j}(t)=\sum_{q=1}^{6} \tilde{G}_{c j q}\left(t, \underline{\boldsymbol{\xi}_{c}}\right) * M_{q}(t)+e_{c j}(t)
$$

with

$$
M_{q}(t)=\iiint_{V} \tilde{M}_{q}(t, \xi) d \xi
$$

Assumption 2: One earthquake has one focal mechanism.

$$
\begin{aligned}
& M_{q}(t)=m_{q} \times \dot{M}_{0}(t) \\
& u_{c j}(t)=\sum_{q=1}^{6} m_{q} \times M_{0}(t) * \tilde{G}_{c i q}\left(t, \boldsymbol{\xi}_{c}\right)+e_{c j}(t)
\end{aligned}
$$

## Applying Low-pass filter

$$
\begin{aligned}
& d_{c j}(t)=F(t) * u_{c j}(t)=\sum_{q=1}^{6} m_{q} \times\left[F(t) * M_{0}(t)\right] * \tilde{G}_{c i q}\left(t, \boldsymbol{\xi}_{c}\right)+e_{c j}(t) \\
& \simeq \sum_{q=1}^{6} m_{q} \times[F(t) * S(t)] * \tilde{G}_{c i q}\left(t, \boldsymbol{\xi}_{c}\right)+e_{c i}(t) \\
& \text { simple function } \\
& \text { (e.q. delta function) } \\
& d_{c j}(t)=\sum_{q=1}^{6} m_{q} \times G_{c i q}\left(t, \boldsymbol{\xi}_{c}\right)+e_{c i}(t) \\
& \text { with } \\
& G_{c i q}\left(t, \boldsymbol{\xi}_{c}\right)=\tilde{G}_{c i q}\left(t, \boldsymbol{\xi}_{c}\right) * F(t)
\end{aligned}
$$

Vector Form:

$$
\mathbf{d}=\mathbf{A}\left(\xi_{c}\right) \mathbf{m}+\mathbf{e}
$$

The solution of the above matrix equation is obtained by least square approach if we assume the centroid location.

$$
\hat{\mathbf{m}}=\left[\mathbf{A}^{T}\left(\xi_{c}\right) \mathbf{A}\left(\xi_{c}\right)\right]^{-1} \mathbf{A}^{T}\left(\xi_{c}\right) \mathbf{d}
$$

We can estimate optimal the centroid using the grid-search method, which minimizes normalized L2-norm as

$$
\begin{aligned}
& \left\|\mathbf{d}-\mathbf{A}\left(\xi_{c}\right) \hat{\mathbf{m}}\right\|_{2} /\|\mathbf{d}\|_{2} \Rightarrow \min \\
& \text { L2-norm: } \quad\|\mathbf{x}\|_{2}=\left(\sum_{n=1}^{N} x_{n}^{2}\right)^{1 / 2}
\end{aligned}
$$

## Moment Tensor to the two Fault Planes

Transformation from moment tensor to the two fault planes.
First we obtain the eigenvectors $(\mathbf{t}, \mathbf{b}, \mathbf{p})$ of moment tensor.

$$
\left(\begin{array}{lll}
M_{11} & M_{12} & M_{13} \\
M_{21} & M_{22} & M_{23} \\
M_{31} & M_{32} & M_{33}
\end{array}\right)=\left(\begin{array}{lll}
\mathbf{t} & \mathbf{b} & \mathbf{p}
\end{array}\right)\left(\begin{array}{ccc}
M_{0} & 0 & 0 \\
0 & 0 & 0 \\
0 & 0 & -M_{0}
\end{array}\right)\left(\begin{array}{lll}
\mathbf{t} & \mathbf{b} & \mathbf{p}
\end{array}\right)^{T}
$$

Nest, we obtain fault vector (n: unit normal vector to fault plane, d: unit slip vector) form the eigenvectors $(\mathbf{t}, \mathbf{b}, \mathbf{p})$ using the equation:

One fault plane model : $\mathbf{n}=\frac{1}{\sqrt{2}}(\mathbf{t}+\mathbf{p}), \mathbf{d}=\frac{1}{\sqrt{2}}(\mathbf{t}-\mathbf{p})$
or
Other fault plane model : $\mathbf{n}=\frac{1}{\sqrt{2}}(\mathbf{t}-\mathbf{p}), \mathbf{d}=\frac{1}{\sqrt{2}}(\mathbf{t}+\mathbf{p})$

## Moment tensor inversion for middle earthquake using local seismic network

Duration of source time function of middle size < 10 sec

If we apply a low pass filter, we can neglect source time function. In my program for local seismic network, we neglect effect of source time function.

Grid Search parameter:
only location of hypocenter



## Moment tensor inversion for middle earthquake using local seismic network

Step I:
Proceeding for near-source data set
filtering, re-sampling

## Step 2:

Calculation of Green's Function
Step 3:
Inversion


## Moment tensor inversion for middle earthquake using local seismic network

If epicenter $=$ horizontal location of centroid of moment release,


